

M<sup>a</sup> Ondina FIGUEIREDO & Teresa PEREIRA da SILVA

LNEG, Geol. Data C., Apt. 7586, 2721-866 Alfragide, & CENIMAT/I3N, Mat. Sci. Dpt., Fac. Sci. Techn., New Univ. Lisbon, 2829-516 Caparica, Portugal

## Introduction

Indium is nowadays widely used in many technologic fields - electronics, solders, low melting-temperature alloys, etc. Discovered in 1863 and isolated four years later as a metal, indium became one of the most relevant scarce metals used in the production of new "high-tech devices" based on innovative nanotechnologies.

Suggestive examples of indium incorporation - along with other rare metals, gallium and tin - are liquid crystal displays (LCDs), organic light emitting diodes (OLEDs) and the recently introduced transparent flexible thin-films (TFTs) [1], manufactured with ionic amorphous oxide semi-conductors (IAOs) within the systems In-Sn-O (ITO) and In-Ga-Zn-O (IGZO). Since the early 1970s, ternary semi-conducting chalcopyrite-type compounds turned into promising materials for photovoltaic solar cells with increased efficiency - e.g., Cu(In,Ga)Se<sub>2</sub> (CIGS).

Consumption of indium is therefore expected to increase in years to come, focusing a special interest on its exploitation from polymetallic sulphide ores [2] - namely, the Iberian Pyrite Belt (IPB), a metallogenic province crossing the south of Portugal and extending to Spain (fig.1) - and on the improvement of indium recovery and recycling technologies. Accordingly, it became opportune to revise indium crystal chemistry, starting with minerals (natural materials).

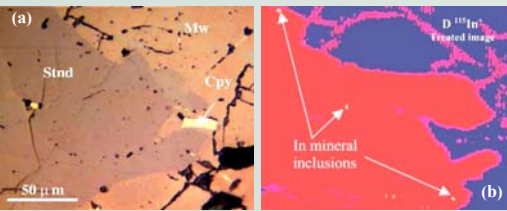
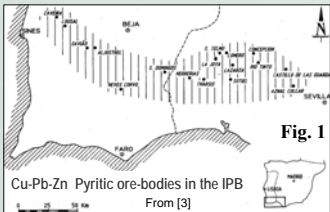


Table 1 - MORPHOTROPIC DOMAINS of CHALCOGENIDE MINERALS (Sulphides & Sulphosalts)

(chemical range for stable diadochic substitutions in minerals deduced from stable synthetic compounds)

Group	I	II	Periodic classification of the Elements										III	IV	V	VI	VI	0		
Period	1	2											3	4	5	6	7	8	9	
1	H												B	C	N	O	F	Ne	2	
2	Li	Be											Al	Si	P	S	Cl	Ar	10	
3	Na	Mg	Transition Metals										Zn	Ga	Ge	As	Se	Br	18	
4	K	Ca	Sc	Ti	V	Cr	Mn	Fe	Co	Ni	Cu	Zn	Ga	Ge	As	Se	Br	18		
5	Rb	Sr	Y	Zr	Nb	Mo	(Tc)	Ru	Rh	Pd	Ag	Cd	In	Sn	Sb	Te	I	Xe		
6	Cs	Ba	*	Hf	Ta	W	Re	Os	Ir	Pt	Au	Hg	Tl	Pb	Bi	Po	(At)	Rn		
7	(Fr)	Ra	**											(112)	(113)	(114)	(115)	(116)	(117)	(118)

[Periodic Table from K. KRAUSKOPF (1967) Introduction to Geochemistry]

Pyrite Tetrahedral structures Galena-plus-sulphosalts / Anions (blende, wurtzite & allied)

## CRYSTAL STRUCTURE-TYPES (STPs)

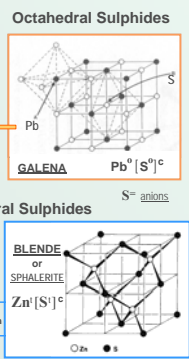
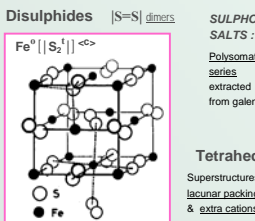


Fig. 2  
Crystal chemical formulae according to J. Lima-de-Faria & M.O. Figueiredo (1976) J. Solid St. Chem. 16: 7-12

## Crystal Chemistry of Indium

Indium (Z=49) has the electronic structure [Kr] 4d<sup>10</sup> 5s<sup>2</sup> 5p<sup>1</sup> and frequently assumes the trivalent state, suggesting the inertness of 5s<sup>2</sup> electron-pair. Like gallium and unlike tin, indium seldom forms specific minerals, exhibiting a distinctly chalcophile behavior in the Earth's crust and occurring rather dispersed within polymetallic sulfide ores (Table 1).

The sulphide roquesite (CuInS<sub>2</sub>) was the first In-mineral to be described [5], followed [6] by indite (FeIn<sub>2</sub>S<sub>4</sub>) and dzhaldinite, a tri-hydroxide with In(OH)<sub>6</sub> octahedra. The recovery of indium stands mostly on the processing of zinc blende - a cubic mineral typifying tetrahedral sulphides (fig.2) where cations fill half of the available tetrahedral sites in a cubic closest packing (ccp) of sulfur anions (S<sup>2-</sup>). The crystal-chemical formula is written Zn<sup>2+</sup>[S<sup>2-</sup>]<sup>c</sup>, where <sup>c</sup> stands for tetrahedral coordination and <sup>h</sup> quotes the anion packing [7]. In is mainly carried in solid solution or diadochic replacement, not only by sphalerite, but also by excess-metal Cu-rich "tetrahedral sulphides" - bornite, Cu<sub>5</sub>Fe<sup>+</sup>[S<sup>2-</sup>]<sup>c</sup> and sakuraiite, (Cu,Ag)<sub>2</sub>(Zn,Fe)(In,Sn)<sup>+</sup>[S<sup>2-</sup>]<sup>c</sup>, plus by complex sulphides of the series tetrahedrite-tennantite also structural derivatives from tetrahedral arrays (fig.3).

Synthetic indium selenides with excess metal were first reported fifty years ago - In<sub>2</sub>Se [9]; since then, many In-chalcogenides have been synthesized - In<sub>6</sub>Se<sub>7</sub>, In<sub>7</sub>Te<sub>10</sub>, InTe and In<sub>4</sub>Te<sub>3</sub> plus In<sub>4</sub>Te<sub>3</sub>, both with excess metal - and their structural characterization has revealed the occurrence of polymetallic In-cations [10,11]: [In<sub>2</sub>]<sup>4+</sup> dimers and/or [In<sub>3</sub>]<sup>5+</sup> trimers (fig.4).

## TETRAHEDRITE-TENNANTITE STP

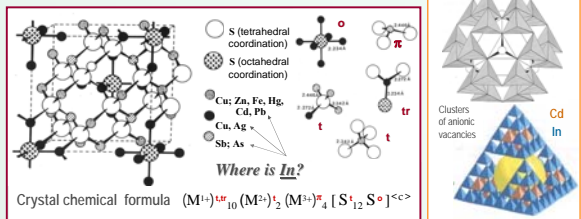


Fig. 4 - Polycations [In<sub>2</sub>]<sup>4+</sup> & [In<sub>3</sub>]<sup>5+</sup> From [10]

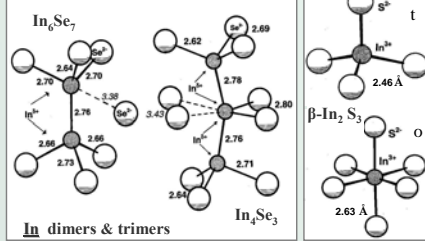


Fig. 6  
Condensed-model standard sheet [14] for A<sub>2</sub>D<sup>o</sup>[X] : large circles, closest packing available interstices.

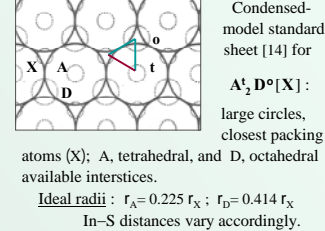


Table 2 - In-concentrations (ppm) in accessory minerals from massive copper sulphide ores of Neves-Corvo polymetallic deposit, southern Portugal [4].

Mineral, empirical formula	In
Bornite, (Cu,Ag) <sub>4.93</sub> Fe <sub>0.97</sub> S <sub>4.10</sub>	not reported
Mawsonite, Cu <sub>3.05</sub> Fe <sub>2.01</sub> Sn <sub>1.0</sub> S <sub>8.05</sub>	300
Tennantite, (Cu,Fe,Zn,Ag) <sub>11.8</sub> (As,Bi) <sub>4.1</sub> S <sub>13.2</sub>	27 400

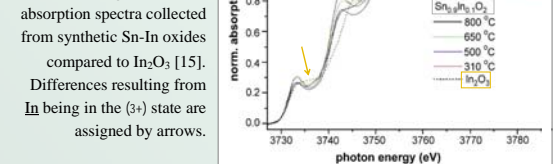
## In at the Nanoscale in Polymetallic Sulphide Ores

Diadochic replacement of indium for zinc is common world-around in sphalerite ores but not so for the IPB polymetallic sulphide ores.

Replacements and recrystallization resulting from a later copper-rich hydrothermal solution often remove the primary texture in volcanogenic massive sulfides (VMS). For that reason, the remobilization of chalcopyrite and tetrahedrite is common in Neves-Corvo copper ores [12] and the resulting paragenesis involves promising indium carriers (Table 2). Encouraging results have already been published [13], showing exsolved indium minerals as minute grains within stannoidite (fig. 5, Table 2).

Mastering the nano-scale complexity in minerals is gaining well turned-out solutions through the use of highly focused micro-beam analytical techniques making use of either particle beams (electrons in EPMA, protons in microPIXE) or photon beams in "third generation" synchrotrons. These large-scale installations provide very suitable instrumental means of non-destructive research at the nanoscale and synchrotron radiation allows for assessing at the binding situation as successfully applied to indium-oxygen compounds.

Fig. 7 - In L<sub>3</sub>-edge X-ray absorption spectra collected from synthetic Sn-In oxides compared to In<sub>2</sub>O<sub>3</sub> [15]. Differences resulting from In being in the (3+) state are assigned by arrows.



## Comments

It is worth noticing that the cubic closest packing anionic array configured by sulphur anions (S<sup>2-</sup>) in chalcopyrite, bornite, mawsonite, stannoidite and even tetrahedrite-tennantite is particularly suitable to accommodate polymetallic cations by filling closely positioned interstitial sites. This is clearly perceptible when looking at a sheet figuring out a single closest packed anionic layer with the available interstices located between this and the forthcoming anionic layer (fig.6): tetrahedral (t) and octahedral (o), in the proportion 2:1. Distances between interstices, [t-o] and [t-t], compare well with In-In distances in synthetic selenides.

Furthermore, once selenides and tellurides are stable synthetic compounds containing In-polycations, it is expectable that Se & Te may be favourably hosted by such anionic closest packing in the presence of indium in natural compounds. Accordingly, it is foreseen to undertake a nanoscale analysis of selected ore samples and pursue an X-ray absorption near-edge spectroscopy (XANES) study at In L<sub>3</sub>-edge (fig. 7) in order to assess In-binding in carrier chalcogenide minerals.

\* The support of the Portuguese Foundation for Science and Technology through project PTDC/CTE-GIN/67027/2006 is acknowledged

References  
[1] E. FORTUNATO, et al. (2005) Adv. Materials 17, 590-594.  
[2] M.O. FIGUEIREDO et al. (2007) Proceedings 9<sup>th</sup> Biennial SGA Mtg., Dublin/Ireland, ed. C. Andrew et al., 1355-1357 (ISBN 0-950989-4-4).  
[3] O.C. GASPARD (1984) Mem. & Notícias, Publ. Museu Lab. Miner. Geol., Univ. Coimbra, 98, 137-150. (in Portuguese).  
[4] O.C. GASPARD (2002) Canad. Miner. 40, 611-636.  
[5] P. PICOT & R. PIERROT (1963) Bull. Soc. Fr. Min. Crist. 86, 7-11.  
[6] A.D. GENKIN & I.V. MURAVEVA (1963) Cf. Amer. Min. 49, 493.  
[7] M.O. FIGUEIREDO & M.J. BASTO (1986) Garcia de Orta, ser. Geol., Lisboa/ICT 9, 41-53 (in Portuguese).  
[8] H. LI, et al. (2003) Angew. Chem. Int. Ed. 42, 1819-1821  
[9] K. SCHUBERT, et al. (1954) Naturwissenschaften 41, 448-453.  
[10] H. SCHWARTZ, et al. (1995) Zeit. Krist. 210, 342-347.  
[11] M. EPPLE, et al. (2000) Zeit. Krist. 215, 445-453.  
[12] O.C. GASPARD & A. PINTO (1991) Miner. Mag. 55, 417-422.  
[13] M. BENZAAZOUA, et al. (2003) Miner. Eng. 16, 1291-1302.  
[14] J. LIMA-DE-FARIA (1994) Structural mineralogy - An introduction. Kluwer Acad. Publ., p.346.  
[15] V. SUBRAMANIAN, et al. (2004) Solid St. Ionics 175, 181-184.